Rivers as Land to Sea Transport Arteries

James P.M. Syvitski Community Surface Dynamic Modeling System UC-Boulder, CO, USA



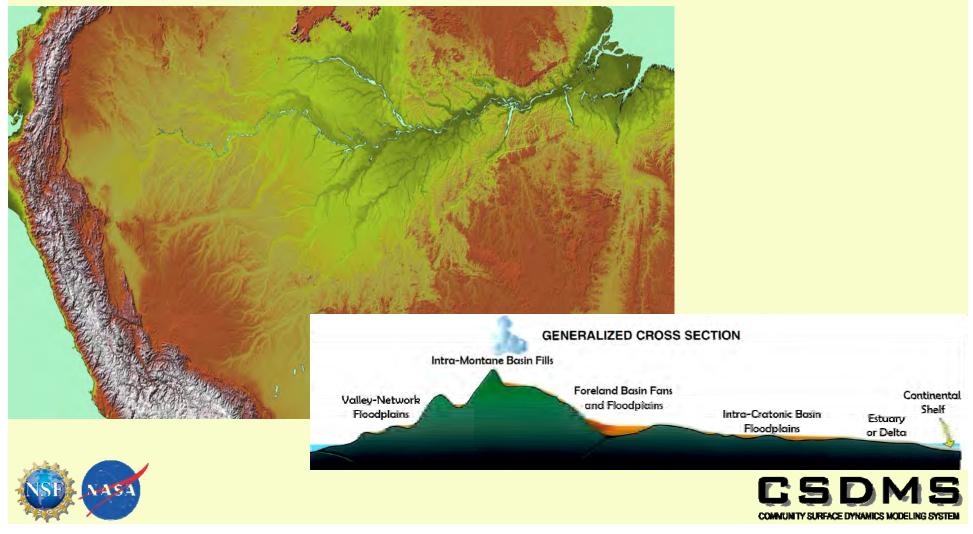






Outline

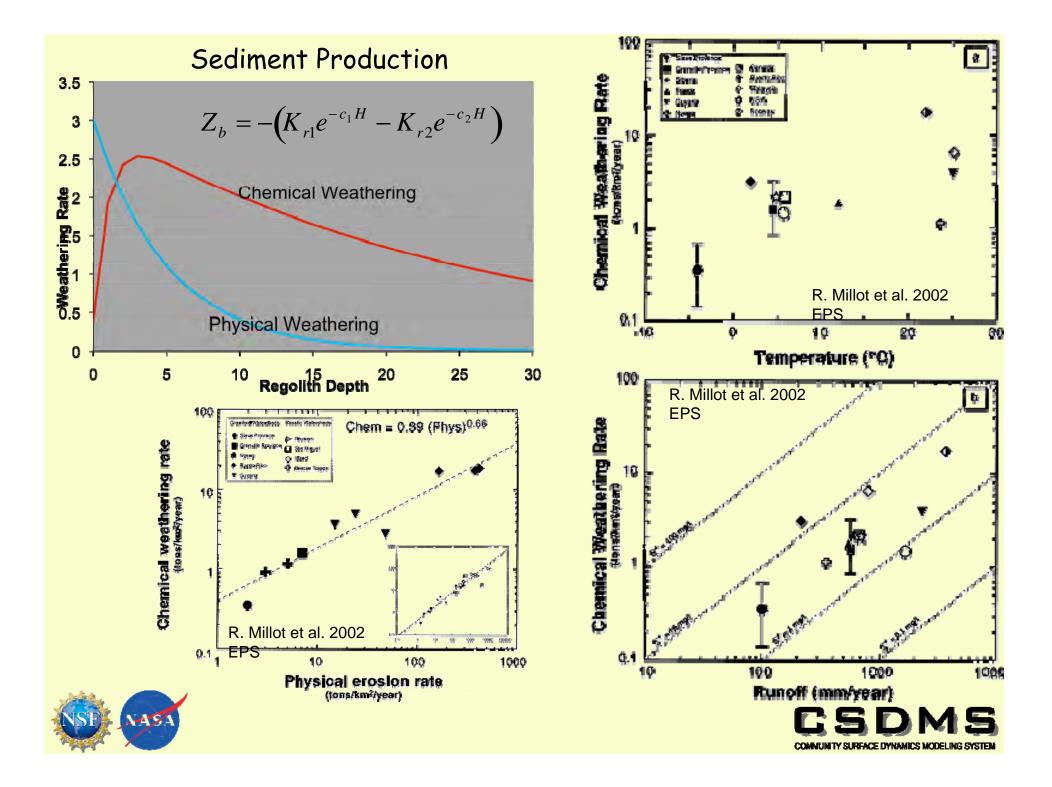
Sediment Production: nature vs. humans Sediment Delivery: bed material load, suspended & wash load Sediment Sequestration: floodplains Sediment Sequestration: deltas

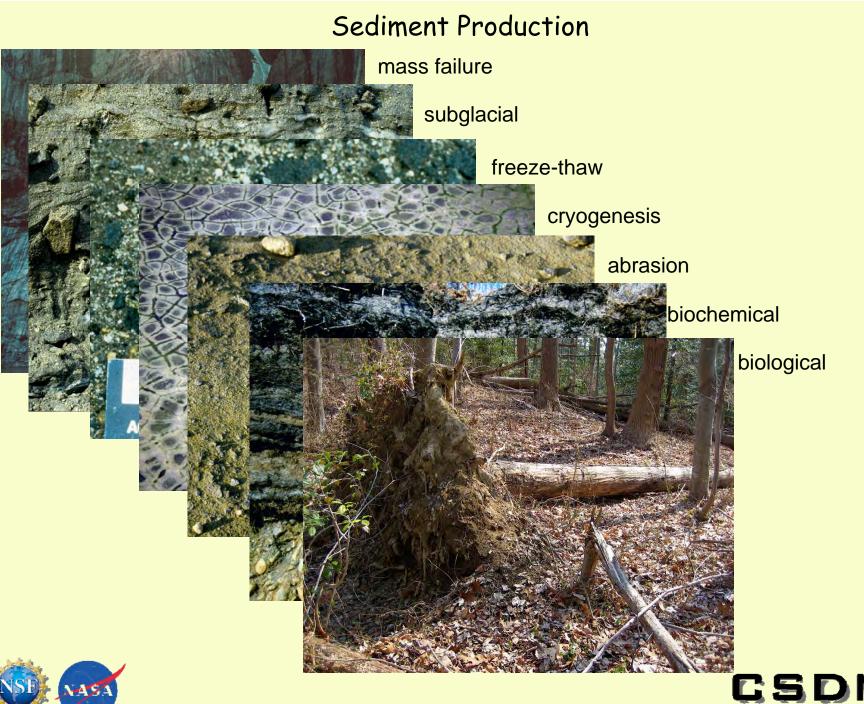


Sediment Production: nature vs. humans









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Sediment Production

SOIL CREEP



SATURATION-EXCESS RUNOFF



THRESHOLD LANDSLIDING



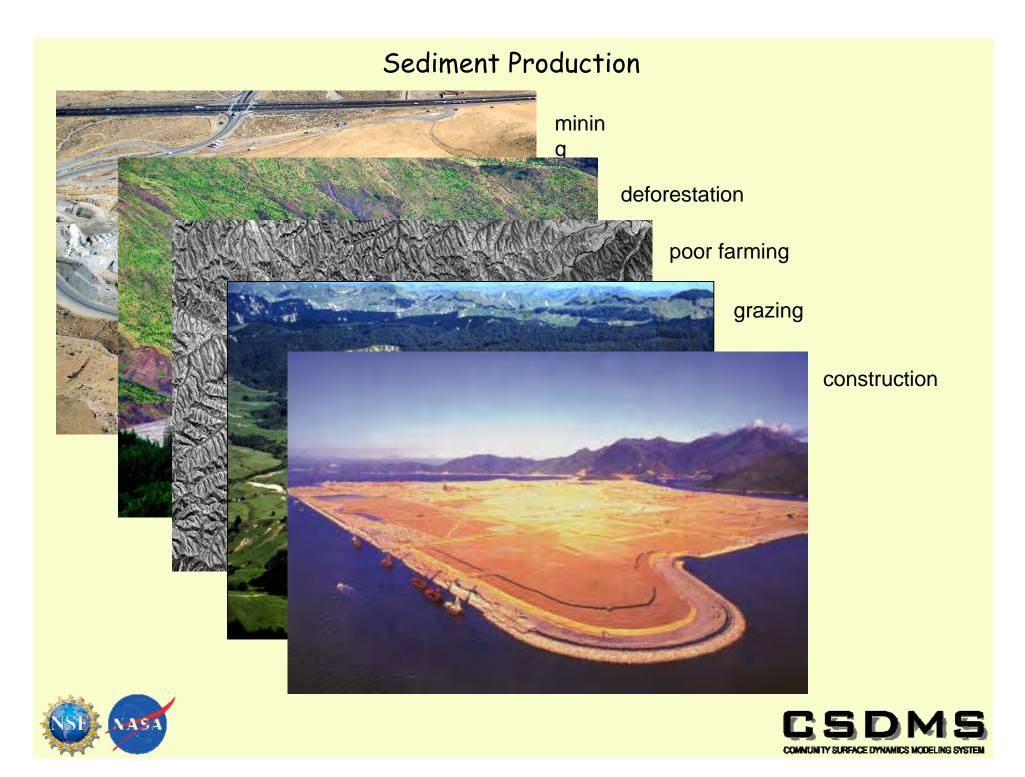
PORE-PRESSURE DRIVEN LANDSLIDING





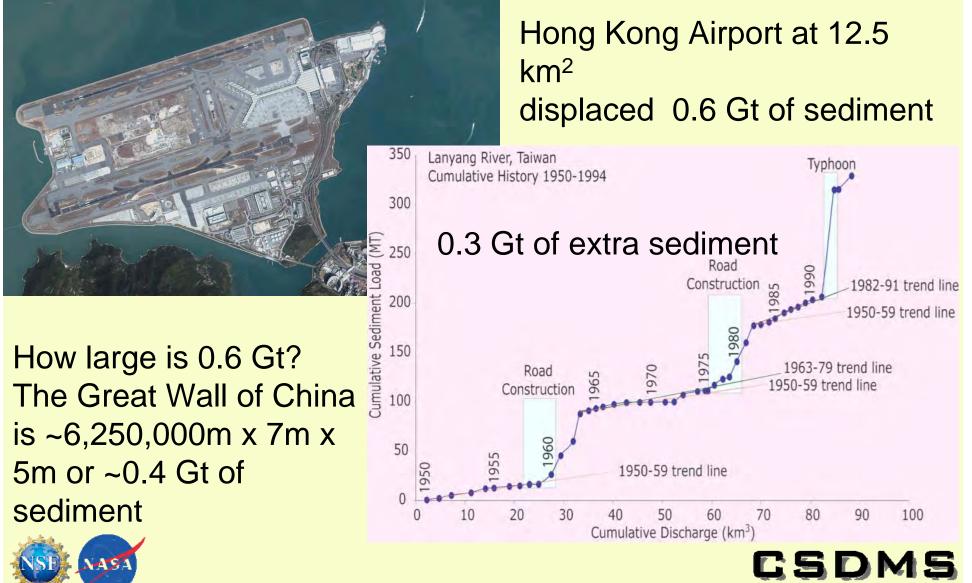
CHILD simulations: G. Tucker, 2002





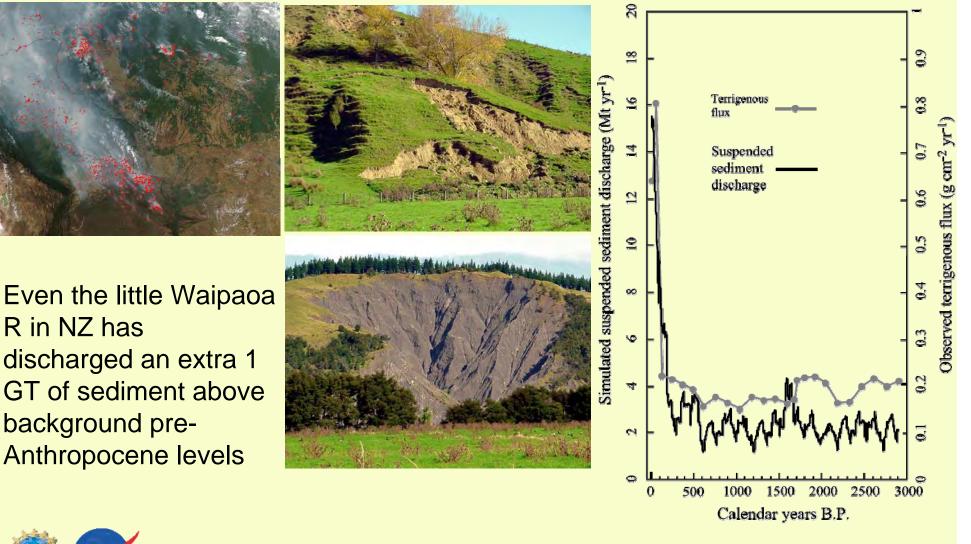
Anthropocene impacts

Transportation systems: gullying, soil erosion, river scouring



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Anthropocene impacts Deforestation: soil erosion, slope failure, and sedimentation;





Anthropocene impacts

Infrastructure and Urbanization: earth surface reshaping



Palm Deira island (not shown) when completed will use 2.0 Gt of sand.





Anthropocene impacts Tillage, terracing: soil erosion, creep, siltation



proliferation of small farms
poor tilling practice
prolonged drought
caused 23.5 million acres to
lose 12.5 Gt of topsoil – Great







Anthropocene impacts

Mining; material displacement, sedimentation, subsidence



Athabaska oil sands, Canada, has 14,000 km² suitable for surface mining. Syncrude mine, one of many, is the largest at 191 km² & processed >30 Gt of sediment.

Hull-Rust-Mahoning Fe Mine, Hibbing, Mn: >1.2 Gt of material removed since 1895. Kiruna Fe Mine, Finland: >1 Gt of material removed since 1900.



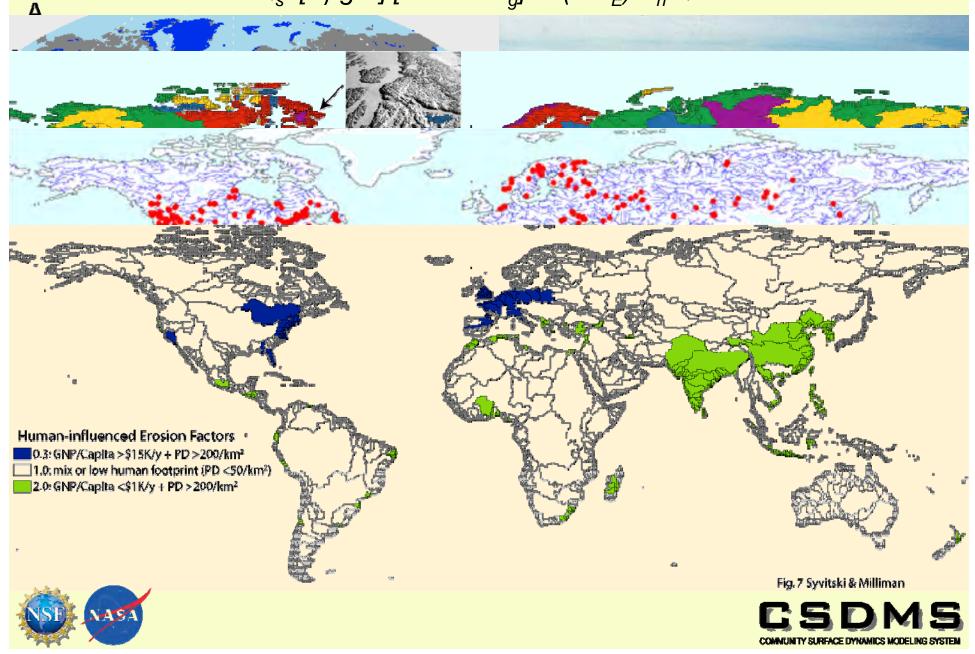


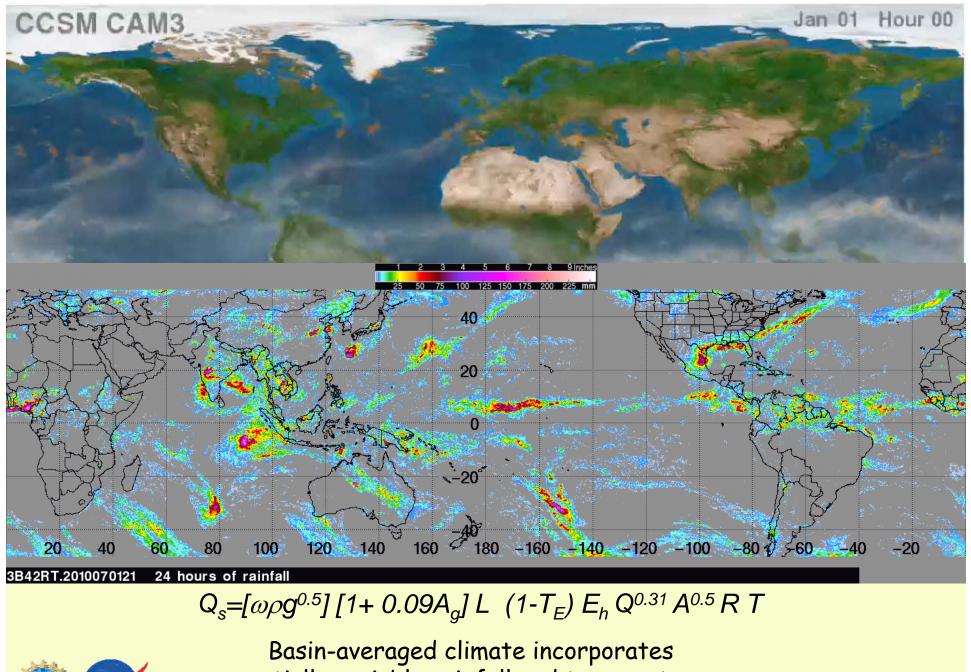
Sediment Delivery: bed material load, suspended & wash load





Sediment Delivery $Q_s = [\omega \rho g^{0.5}] [1 + 0.09A_g] L (1-T_E) E_h Q^{0.31} A^{0.5} R T$

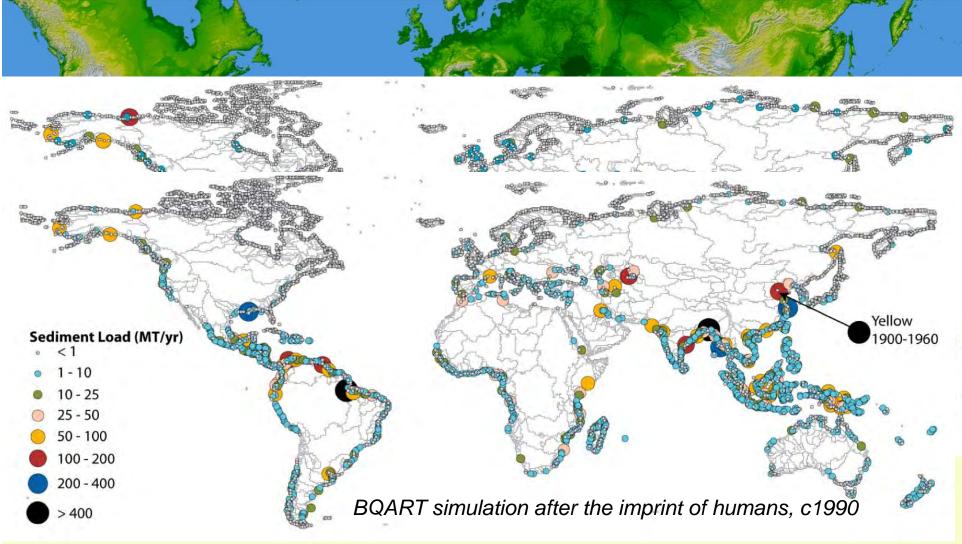




spatially variable rainfall and temperature



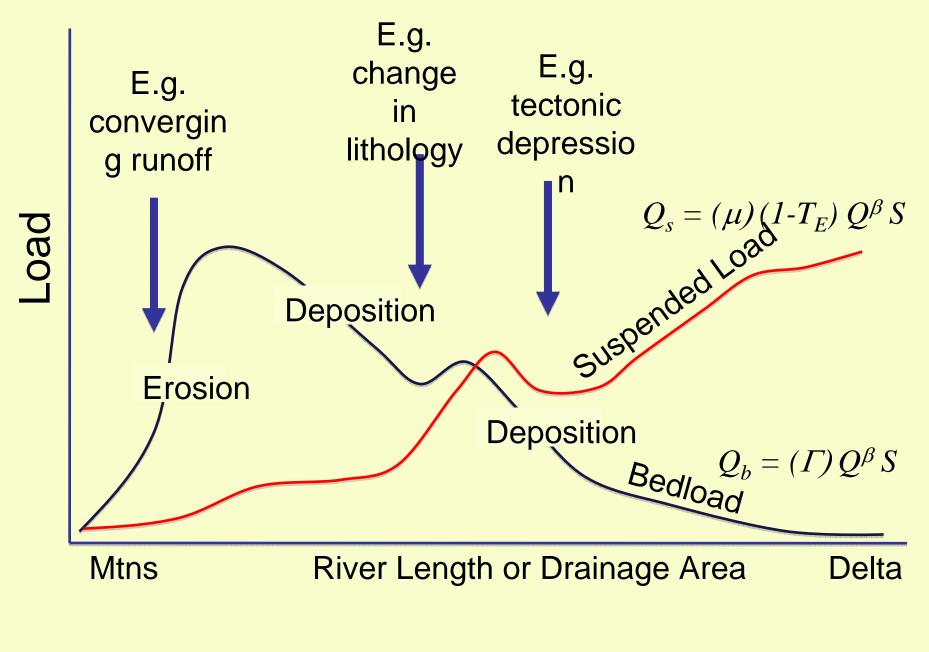
Sediment Delivery $Q_s = [\omega \rho g^{0.5}] [1 + 0.09A_g] L (1-T_E) E_h Q^{0.31} A^{0.5} R T$





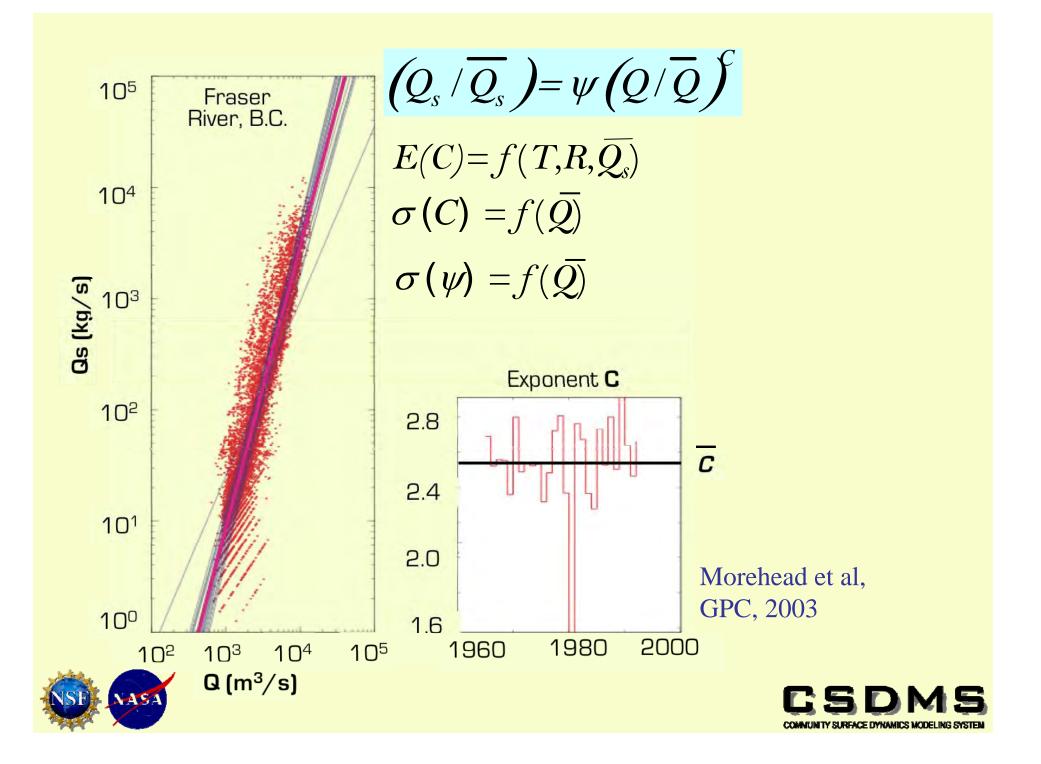
BQART estimates fall on average within 38% of the measured loads on 488 global rivers that drain 63 per cent of the global land surface.

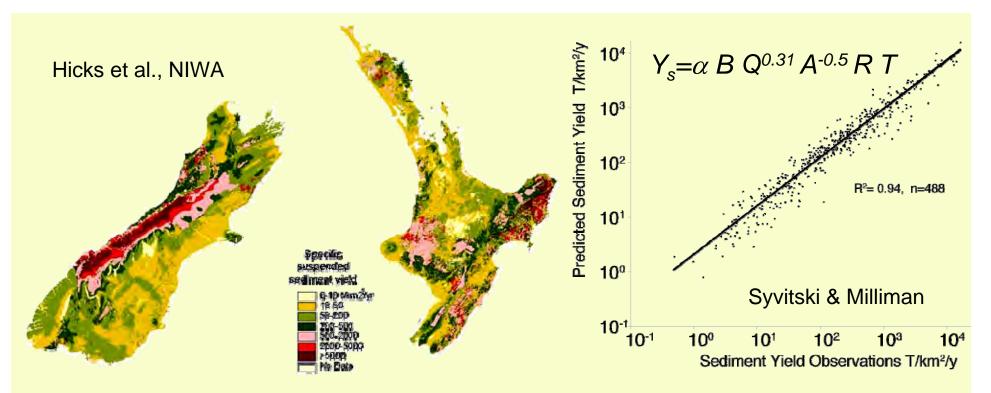


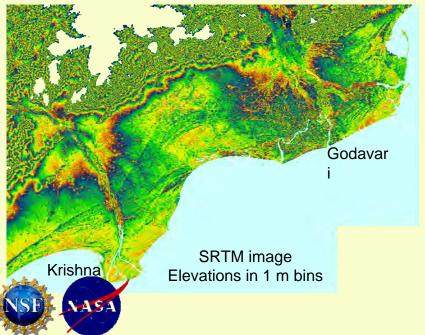








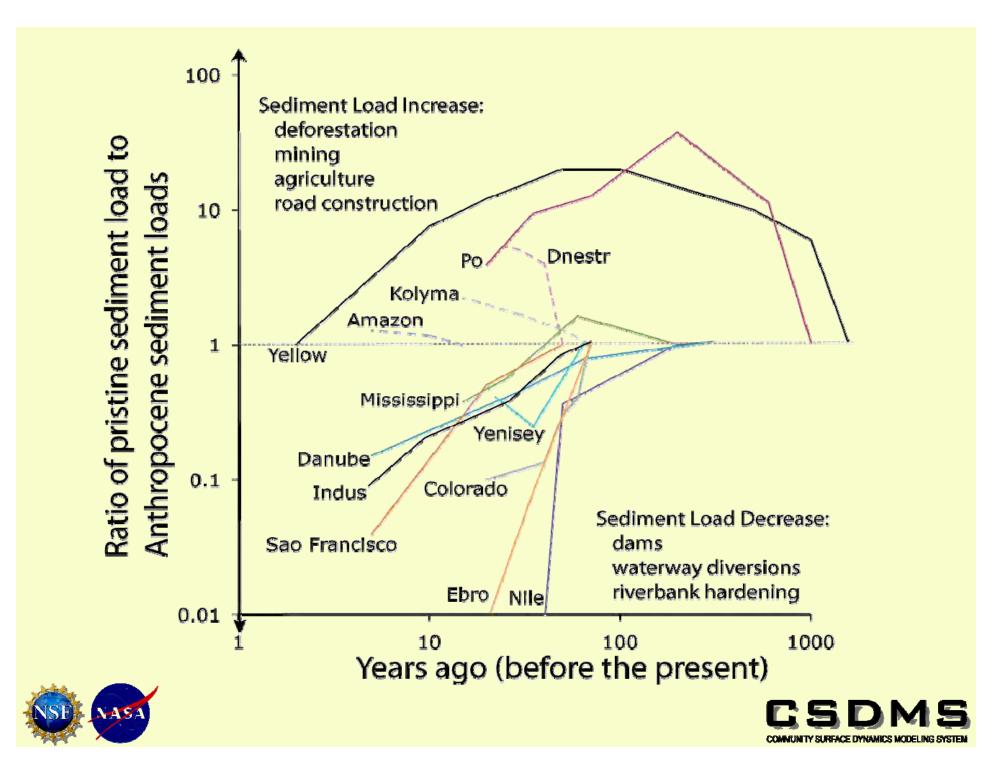




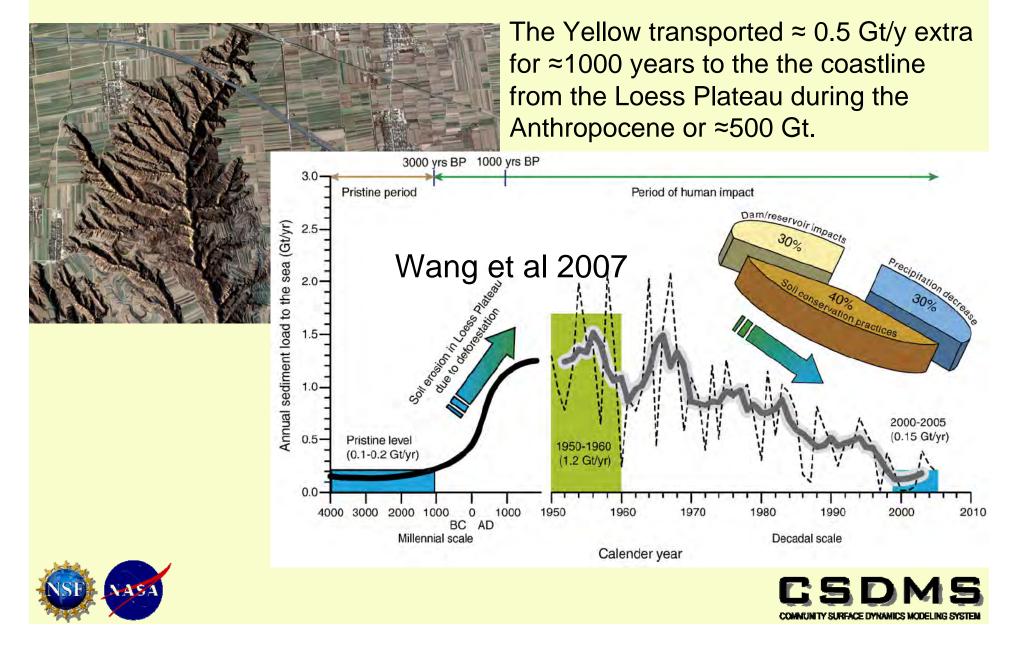
Sediment yield decreases away from highlands because:

 1) Highland sediment delivery is partly trapped on floodplains & delta plains
 2) Local sediment production is low, e.g. low locale relief, rain shadows, vegetation cover





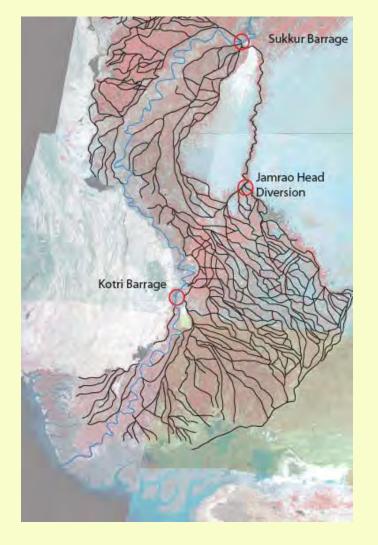
Anthropocene impacts on geomorphology & sediment flux Tillage, terracing: soil erosion, creep, siltation



Anthropocene impacts

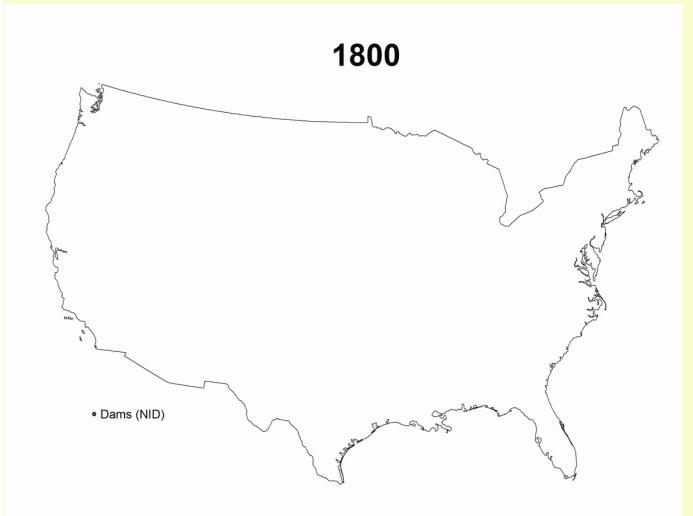
Waterway re-plumbing: reservoirs and dams, diversions, channel levees, channel deepening, discharge focusing, and ultimately coastline erosion;









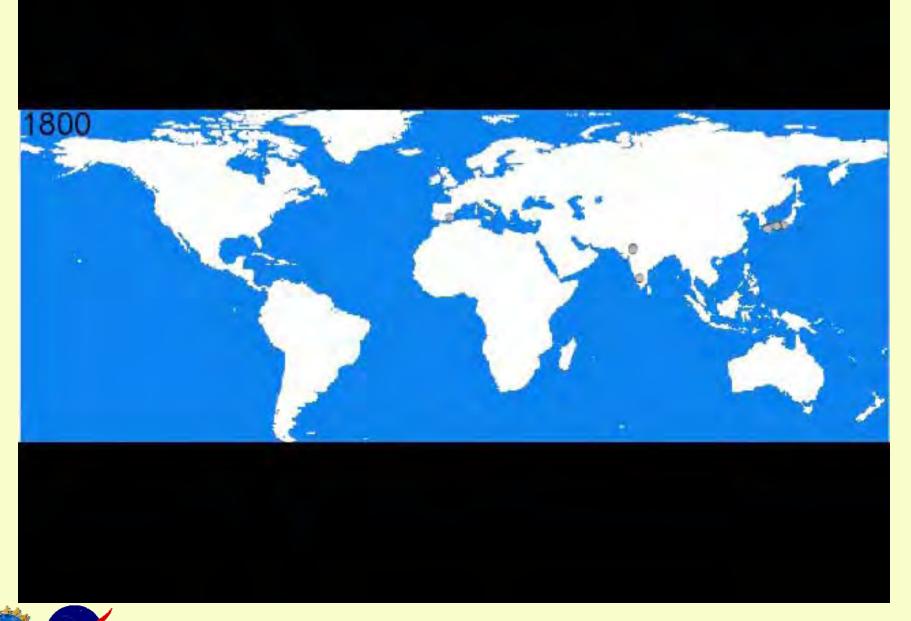


Reservoirs

2.3 Gt/y LESS sediment reaches the coast worldwide











Continent	<i>Q</i> (km ³ /y)	<i>Qs</i> with Humans (Gt/y)	<i>Qs</i> Pre-Humans (Gt/y)	Sediment Flux as a percent of Total Global Flux Q 2 4 6 8 10 12 80
Africa	3797	1.1	1.6	70 60
Asia	9806	4.8	5.3	50
Australasia	608	0.28	0.24	40 Yellaw,
Europe	2680	0.4	0.6	30 Yangize
Indonesia	4251	2.4	2.4	20 Ganges 10 Ganges
North Americ	a 5819	1.5	1.7	0 Amazon
Oceans	20	0.004	0.003	
South Americ	a 11529	2.4	3.3	-20 MAM%
Global	38510	12.8	15.1	-30 -40

Seasonal Sediment Flux as a percent of the Annual FLux

Modern seasonal sediment load for global rivers: **Green - March-May; Red - June-August; Orange - September-November; Blue - December-October**. Superimposed is the annual sediment flux averaged across 1° of latitude. Major rivers are hot spots of sediment discharge.





Sediment Transfer & Storage: Floodplains



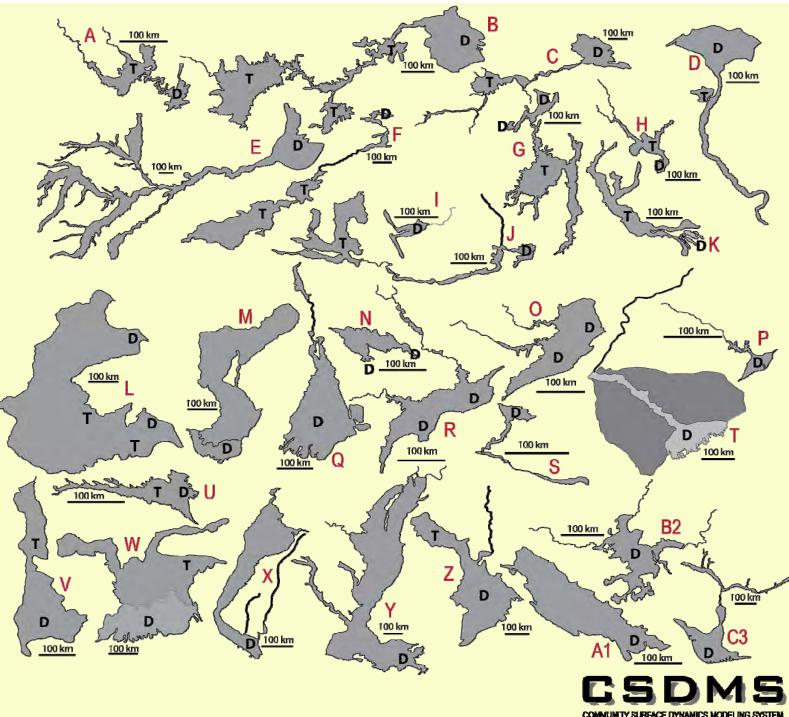


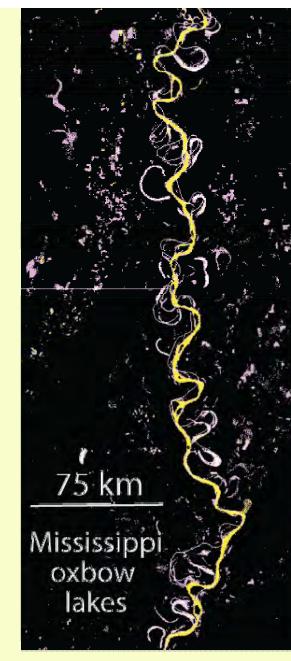


•33 Late Quaternary floodplains & deltas •boundaries are ≤100 m asl

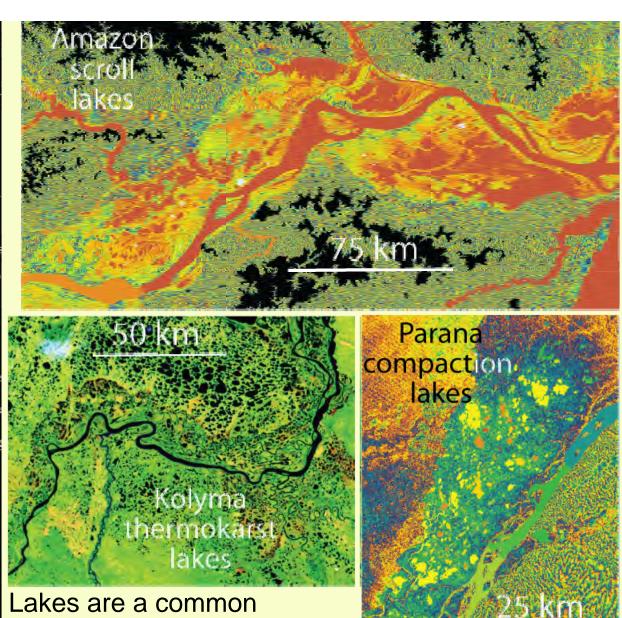
- •**D** = delta;
- •T = tectonic depression









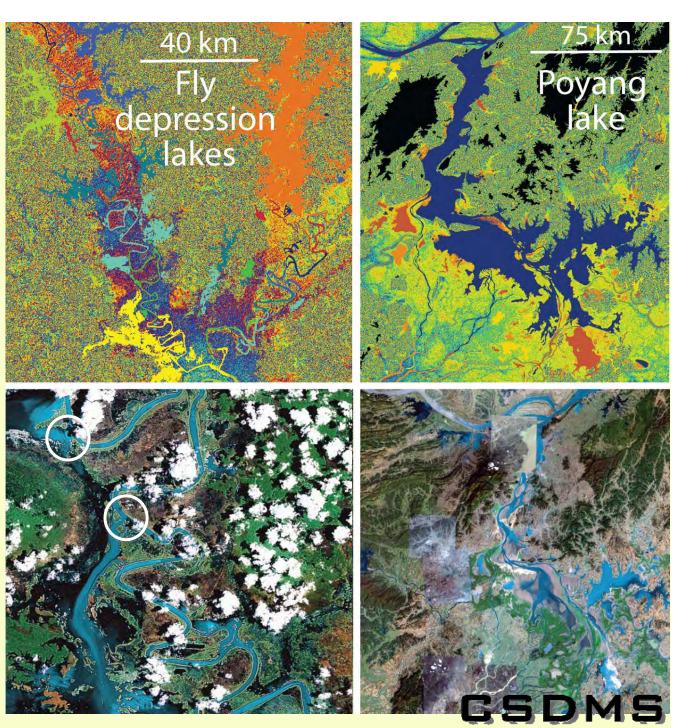


Lakes are a common element of floodplains where flood waters can sequester sediment

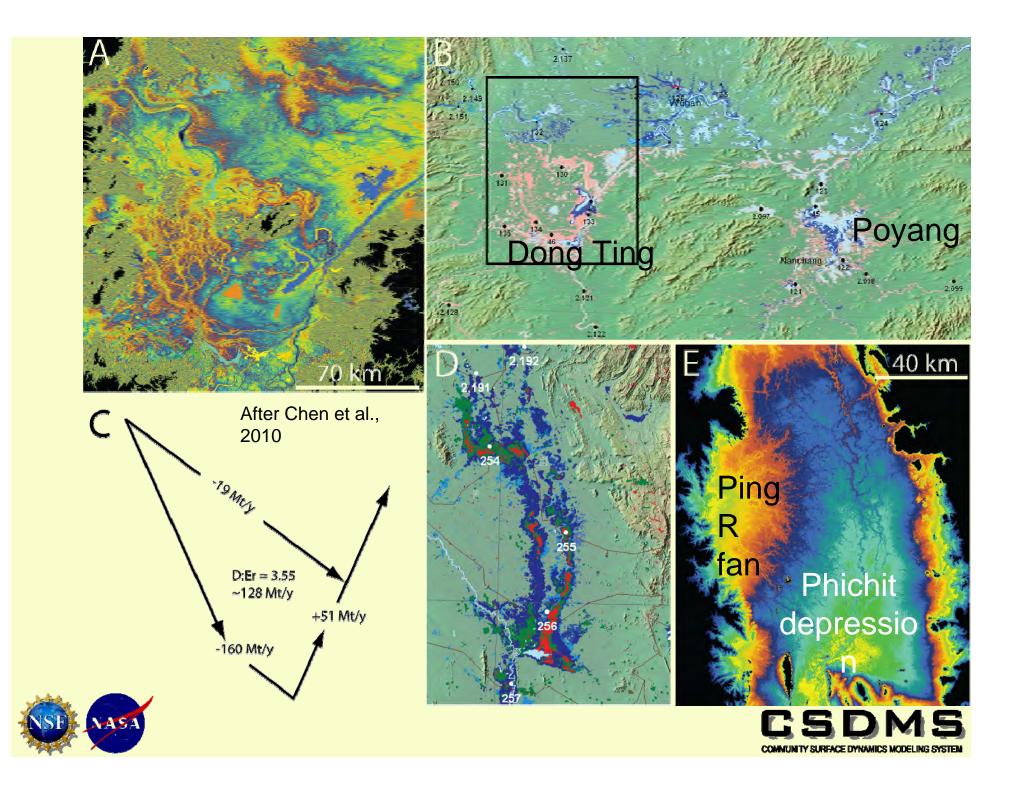


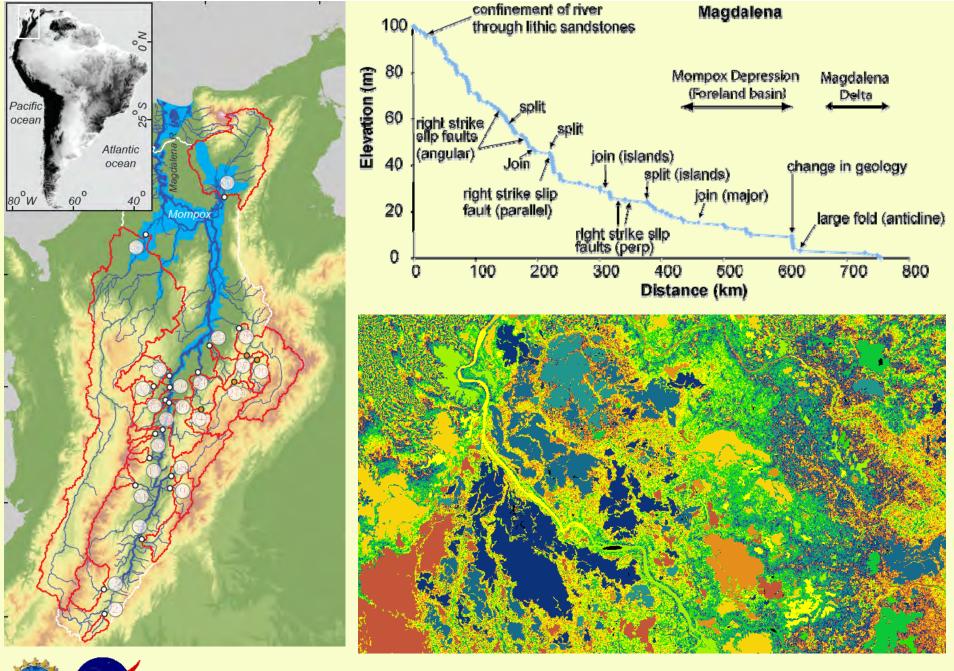
<u>Tectonic</u> <u>Depressions:</u> have a few common characteristics — no one is diagnostic: down-valley slopes;

- expanded valley widths;
- multiple overflow channels similar to deltaic distributary channels;
- Iakes connected to the main river channel;
- highly prone to flooding, often annually;
 highly prone to flooding, often
 annually;
 bighly prone to flooding, often
 annually;
 bighly prone to flooding, often



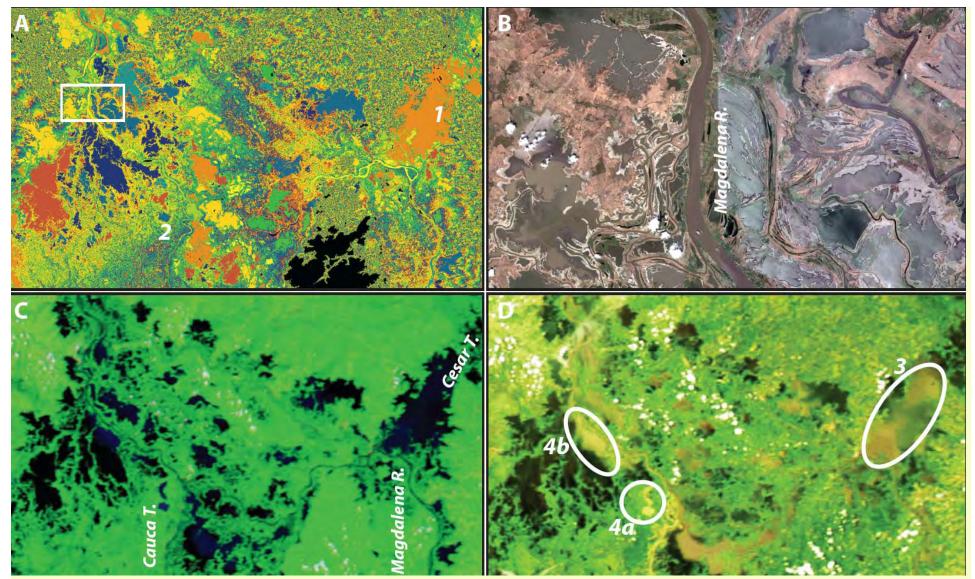
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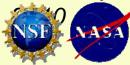




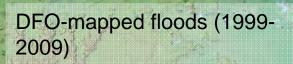




Mompox depression: overflow & levee failures cause extensive flooding (April – Nov) Over last 7500 y, aggradation rate is 3–4 mm/y with deposits 10m to 130m thick With 27 yr of observation, 14% of Magdalena sediment load is trapped *Kettner et al.,*



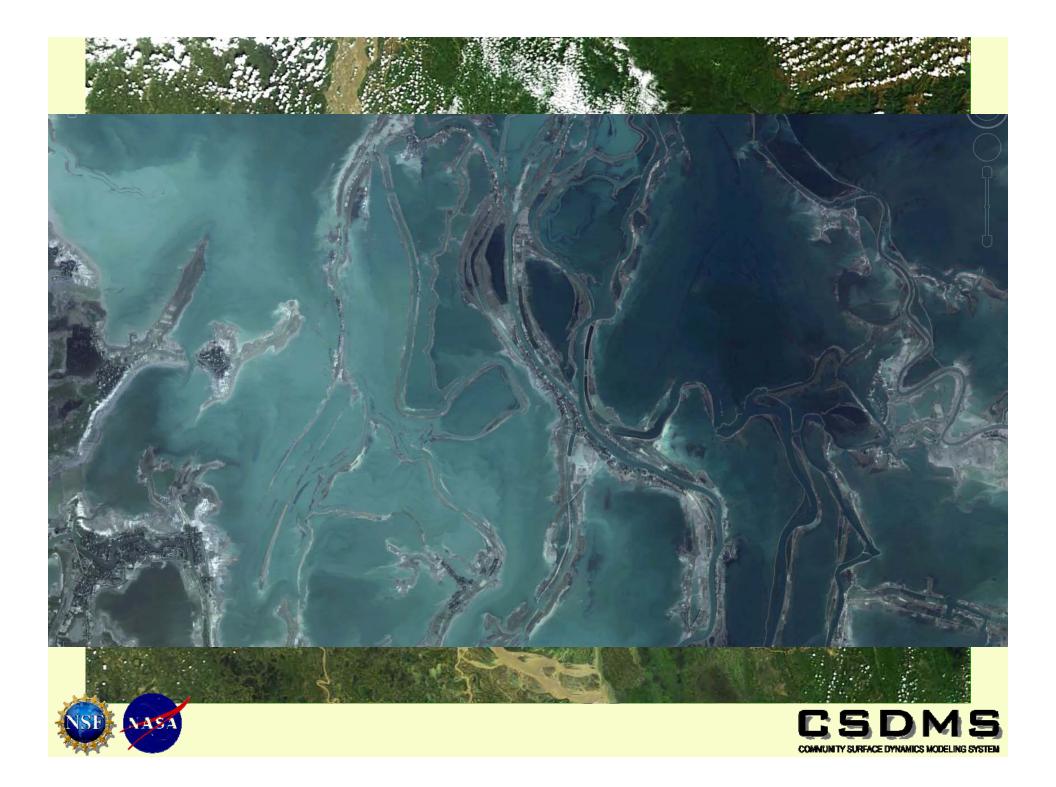


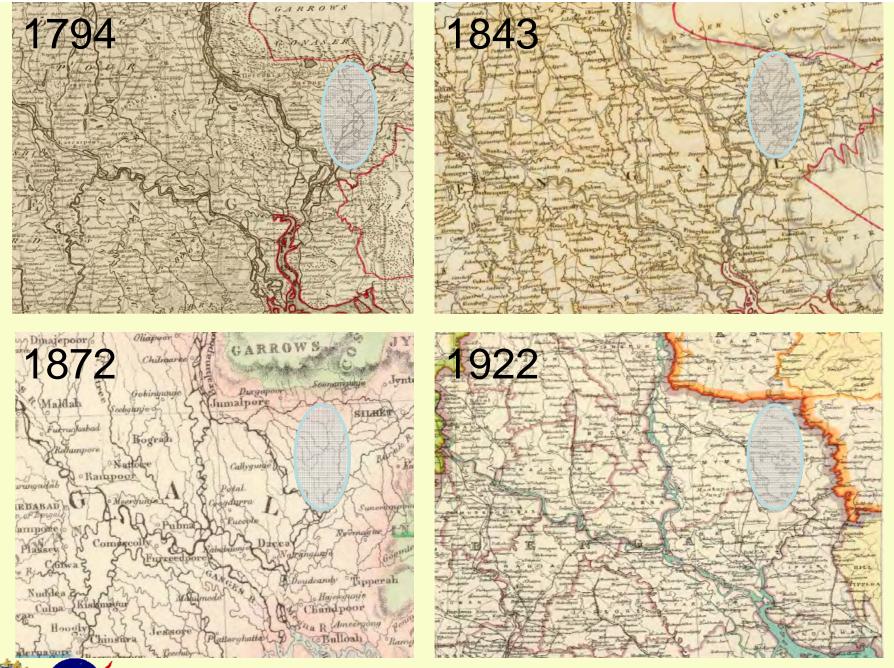


Recent flooded areas overlie earlier flooded areas



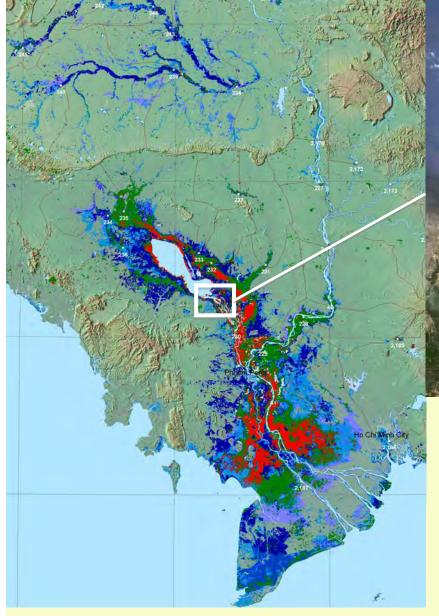










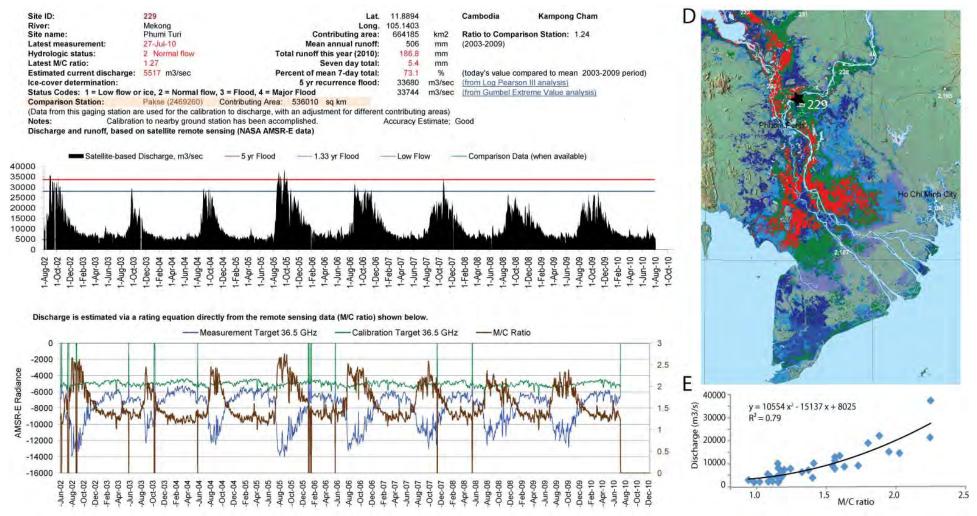




Mekong River (Vietnam) flooding (backfilling) of the tectonic depression Tonle Sap (Cambodia) with both water and sediment





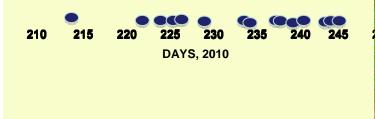


Microwave-sensed river discharge: A) DFO-River Watch station 229 July 27, 2010, Mekong River. B) Daily discharge-AMSRE (2002-2010). C) Daily radiance river reach target, calibration target & M/C ratio. D) Station location. E) International station calibration employs monthly means where station data are available. U.S. and European River calibration uses daily in situ discharge data.



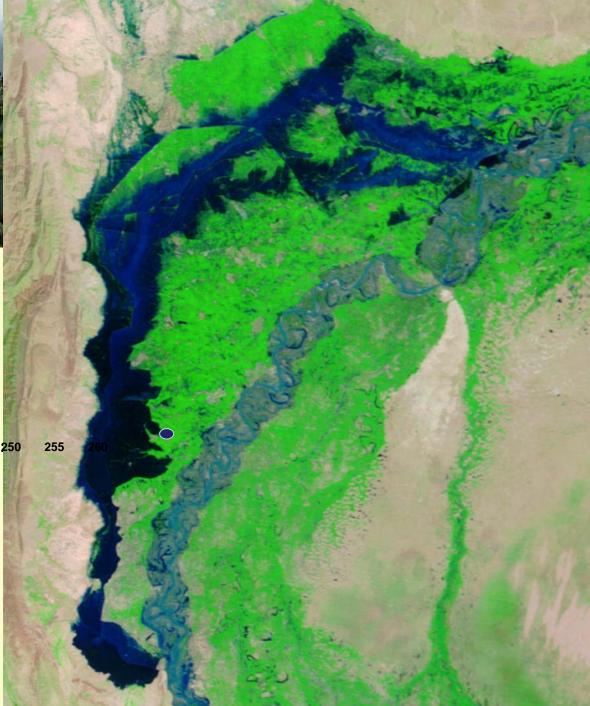


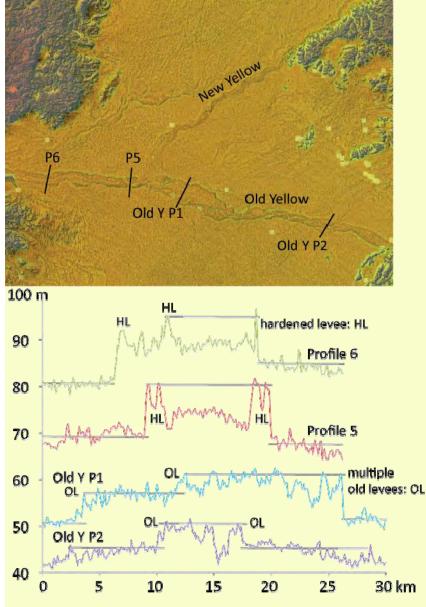
Images are GeoTiffs from either MODIS-Terra or MODIS-Aqua

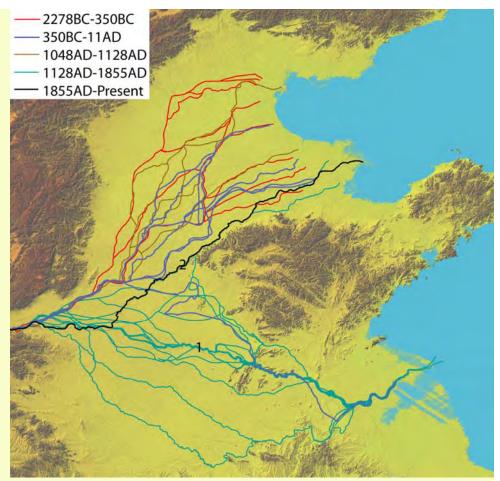


100 km









The Yellow once freely migrated across its 700 km wide floodplain and is now fixed as a narrow elevated floodplain.

The elevated deposit is 50 Gt or 10% of the 500 Gt extra sediment transported by the Yellow to the the coastline.





Sediment Transfer & Storage: Deltas





Controls on Delta Elevation

$$\Delta_{RSL} = A - \Delta E - C_n - C_A \pm M$$

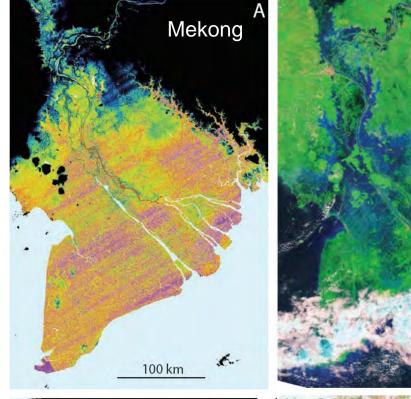
 Δ_{RSL} = Vertical change in delta surface elevation (m/yr)

A = Sediment Aggradation Rate	e (m/yr)
-------------------------------	----------

- ΔE = Eustatic Sea Level Rise (m/yr)
- Cn = Natural Compaction (m/yr)
- Ca = Accelerated Compaction (m/yr)
- M = Crustal Vertical Movement (m/yr)







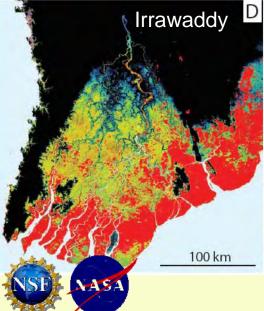
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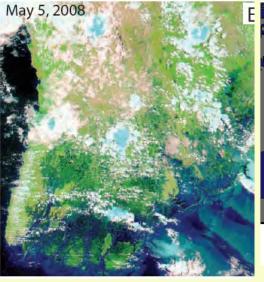
$$\Delta_{RSL} = \pm A - \Delta E - C_n - C_A \pm M$$

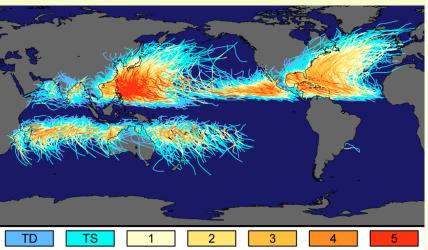
Flooding on deltas can occur from

- 1. Channel overbanking,
- 2. Ocean surges (cyclones, tsunamis)

Aggradation will depend on the sediment flux being carried by the flood waters, and the retention rate on the delta







Saffir-Simpson Hurricane Intensity Scale

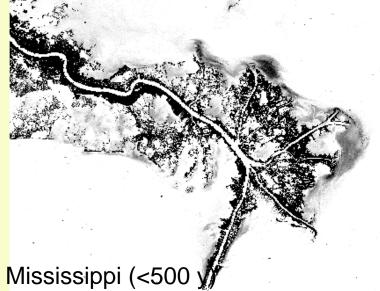


In the pre-dam period when sediment erosion was rampant (1600 – 1950) deltas rapidly expanded, and in some cases formed prograding from coastal plain systems.





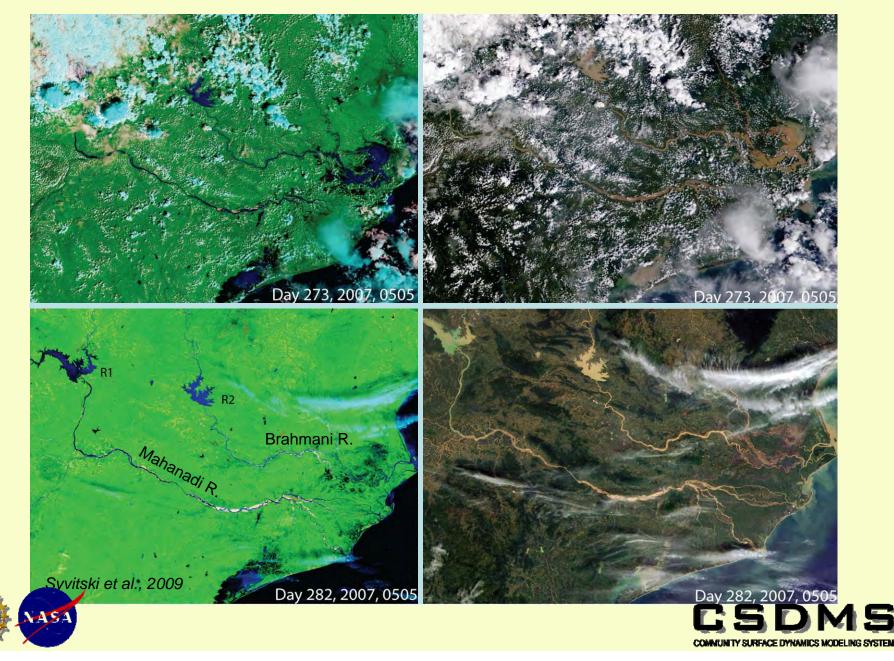






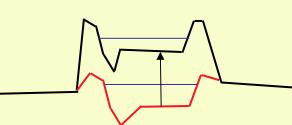


Sedimentation occurred between distributary channels from overbank flooding.

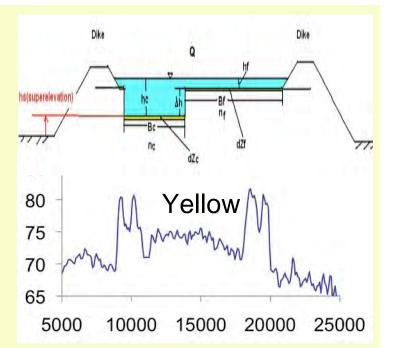


<u>Channel Retention:</u> — Deposition within distributary channels

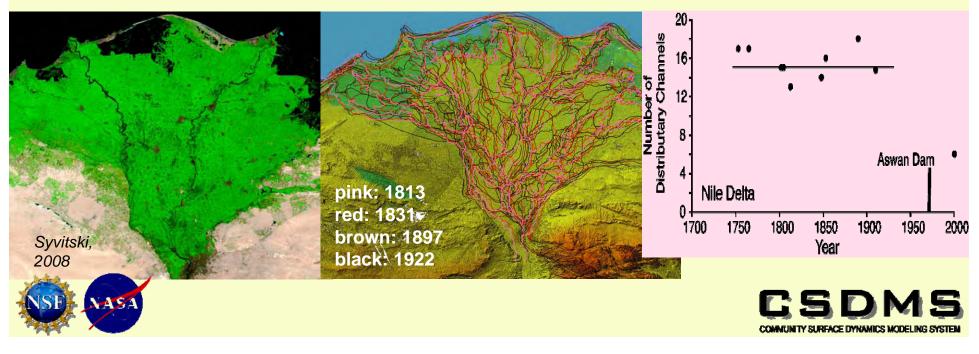




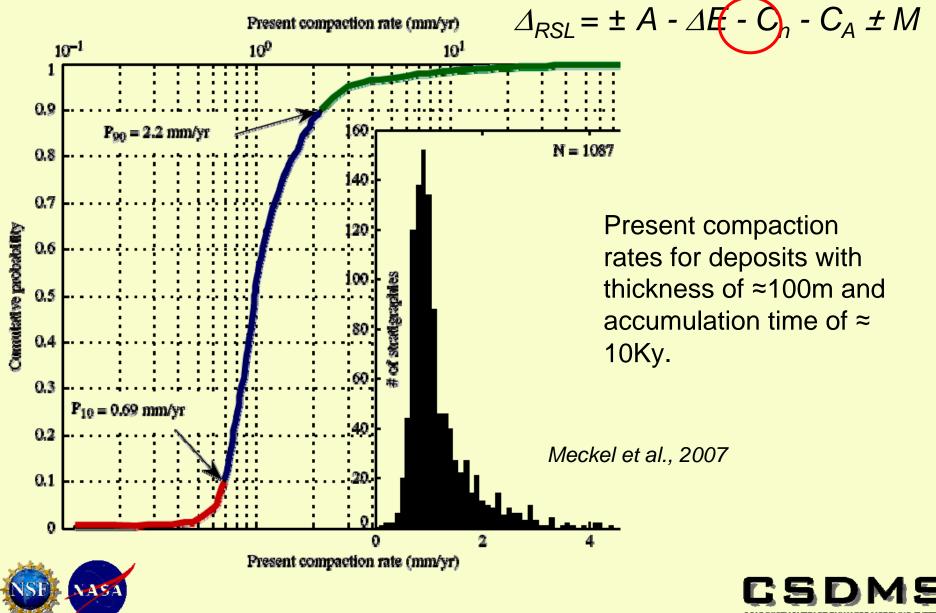
Stop-banks super-elevate the riverbed above the floodplain.



Aggradation from migration of channels.



<u>Natural Compaction Rates</u> changes in the void space within sedimentary layers (dewatering, grain-packing realignment, organic matter oxidation)



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Accelerated Compaction Rates

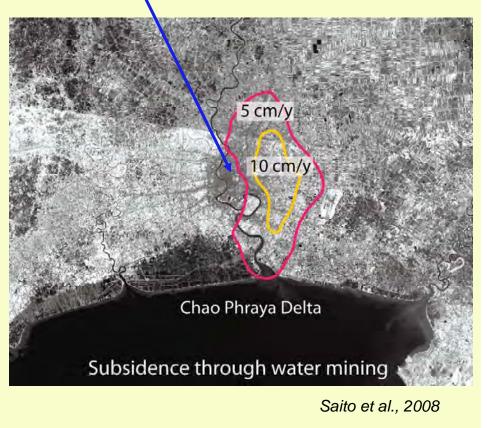
 $\Delta_{RSL} = \pm A - \Delta E - Q_n - Q_A \pm M$



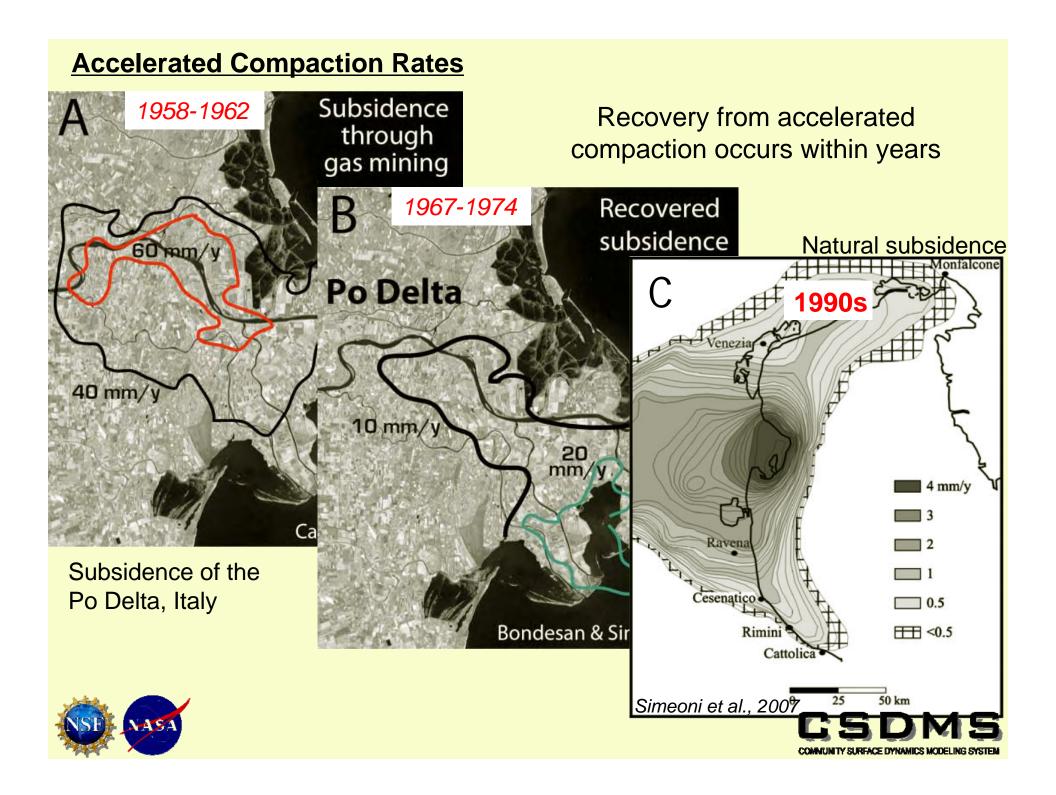


ExamplesYangtze:28 mm/y before controlsNiger:25 to 125 mm/yChao Phraya:50 to 150 mm/yPo:60 mm/y before controls

Bangkok's population went from 1M to 12M in 35 years







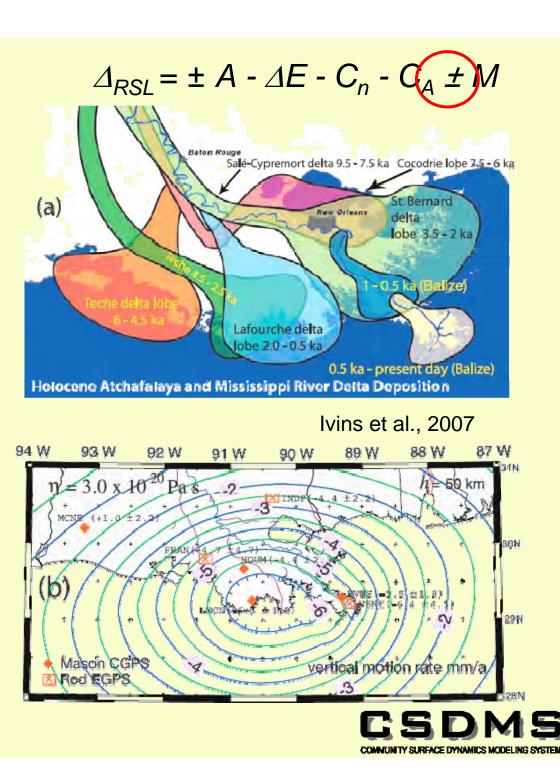
Crustal Subsidence

Each location on a large delta sinks at different rates, depending on their load history.

Mississippi delta lobes weigh between 200 to 900 billion tonnes. Today the various Mississippi lobes are sinking at between:

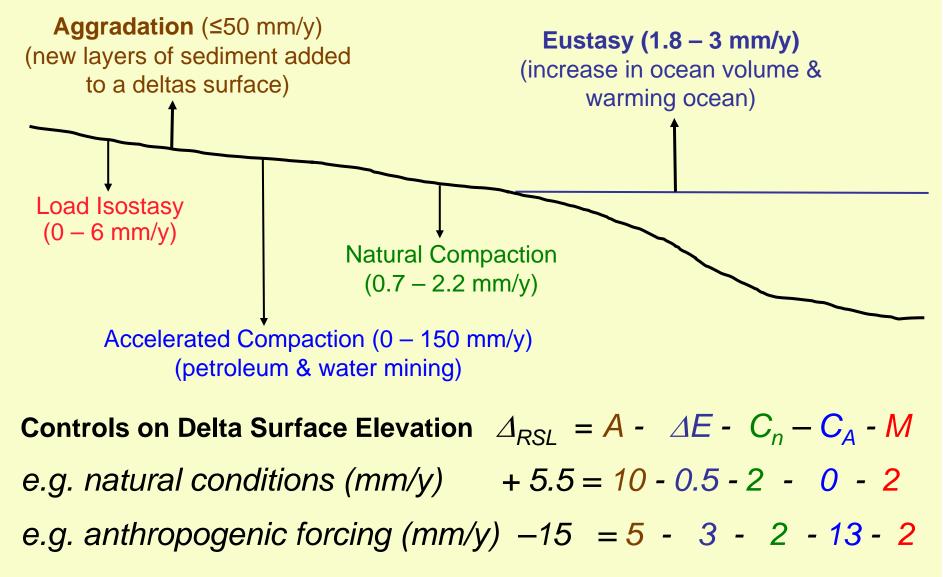
1)0.3 to 3.6 mm/y (Hutton & Syvitski, 2008)

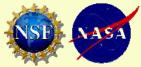
2) 2.0 to 6 mm/y (lvins et al., 2007)



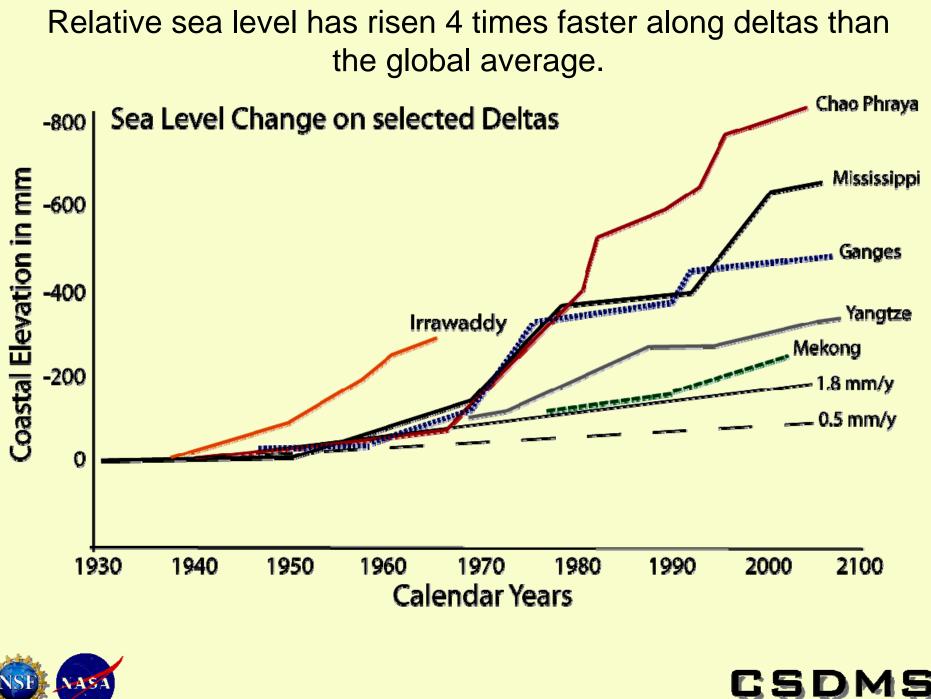


Net Changes in a Delta's Relative Sea Level

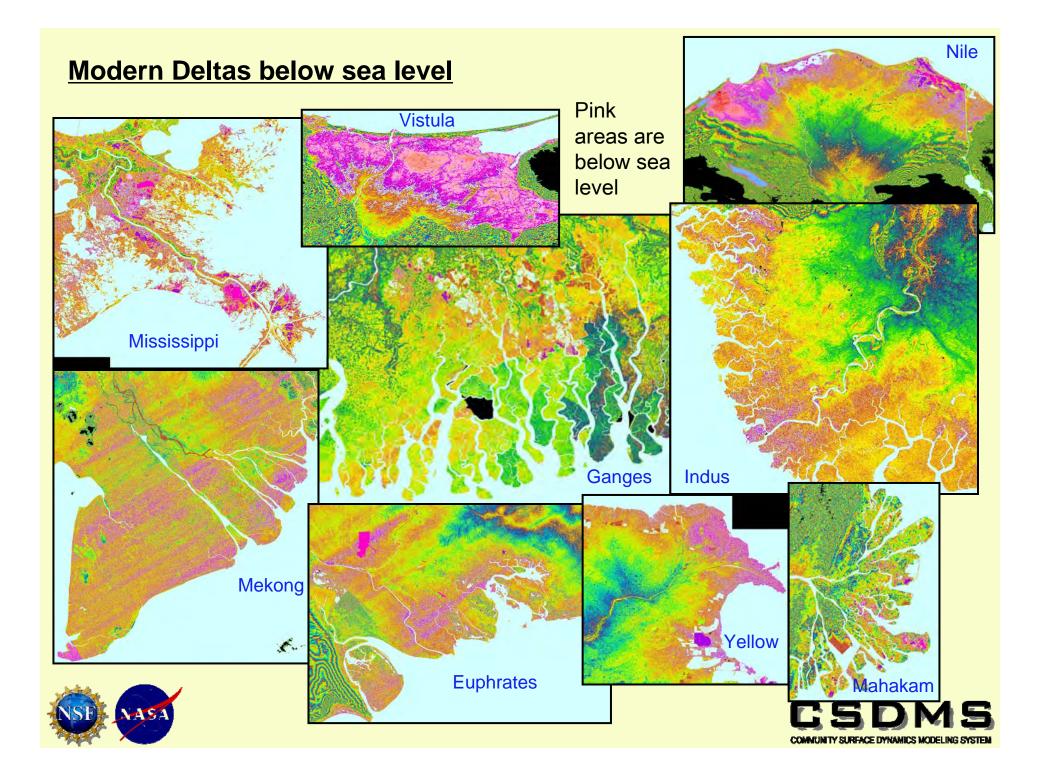




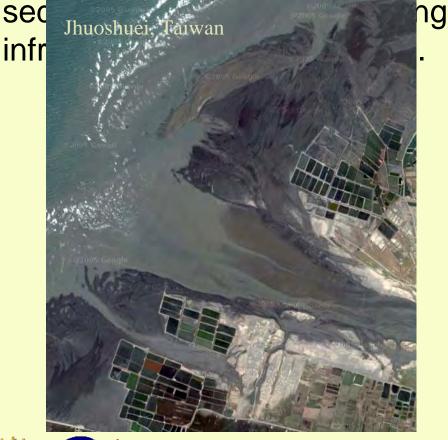


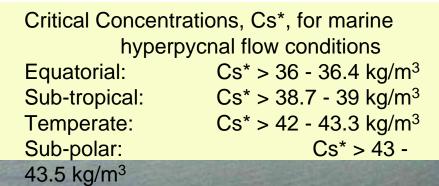


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<u>Hyperpycnal discharge</u> is limited to small & mediumsized rivers that drain mountains capable of generating hyper-elevated





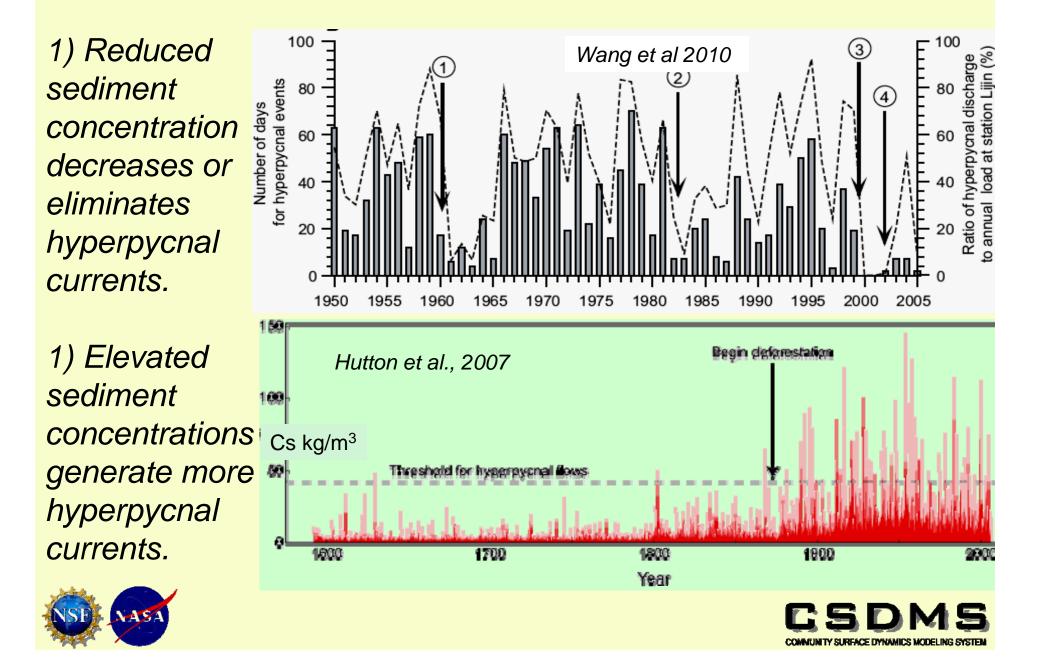
Skeidarasundar, Iceland, 1996

Skeidarasundar, Iceland, 1996





Anthropocene impacts on coastal sediment flux



Summary

➢ We have predictive but basic understanding of sediment production, delivery and storage at global, regional and local scales.

➤ The human footprint is ubiquitous and significant: by 1600 AD soil disturbance rampant; by 1900 AD mechanization, mining, terracing, deforestation lead to global land-sea signals; by 1930 AD subsidence began for many deltas; by 1950 AD sediment sequestration behind dams is a dominant signal in most rivers.

➤ Tectonic depressions located on many (55%) floodplains within ≤ 100 m asl are natural traps of the delivery of sediment to the coastal zone — sequestration mechanisms are highly varied.





➢ global deltas have large areas (>>100,000 km²) <2m a.s.l; most (75%) experienced flooding in the last decade, submerging >260,000 km² of land. Vulnerable low-lying lands are expanding rapidly, due to sinking.

➢ Deltas are sinking 4 times more rapidly than ocean level is rising due human interference in river basins and their deltas due to

1.Reduced sediment delivery to the deltas (>2.3 Gt less sediment reaches these deltas per year)

2. Sediment delivery bypassing the delta plains (fixed distributaries with stop banks).

3.Accelerated compaction due to subterranean mining: 70% of deltas

>Infrastructures for growing mega-cities is a dominant factor.

➢Sediment dispersal into the coastal ocean is strongly influenced by humans



